

Progress on a Common AMSR-E/TMI Level-3 Oceanic Rainfall Algorithm

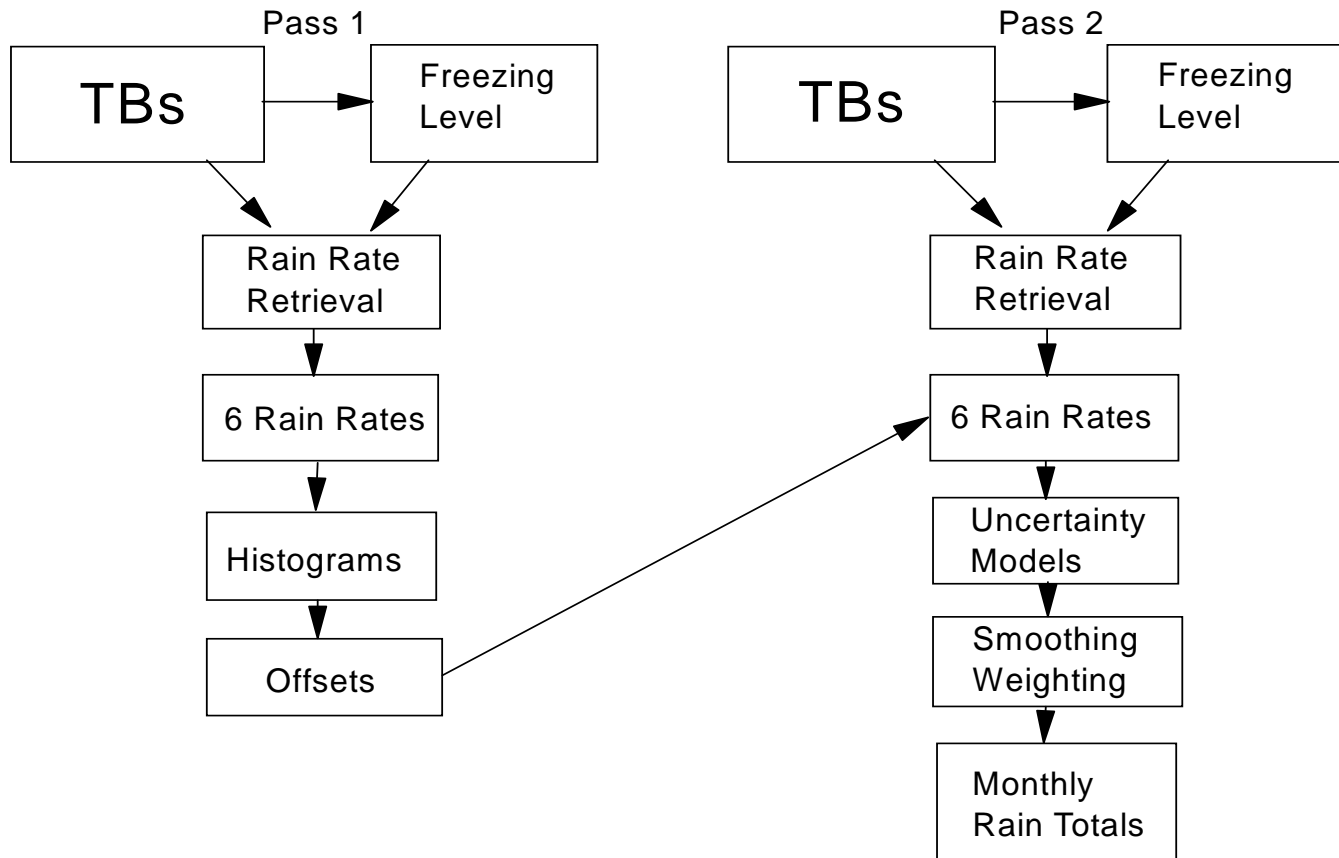
Thomas Wilhelm & Richard Weitz

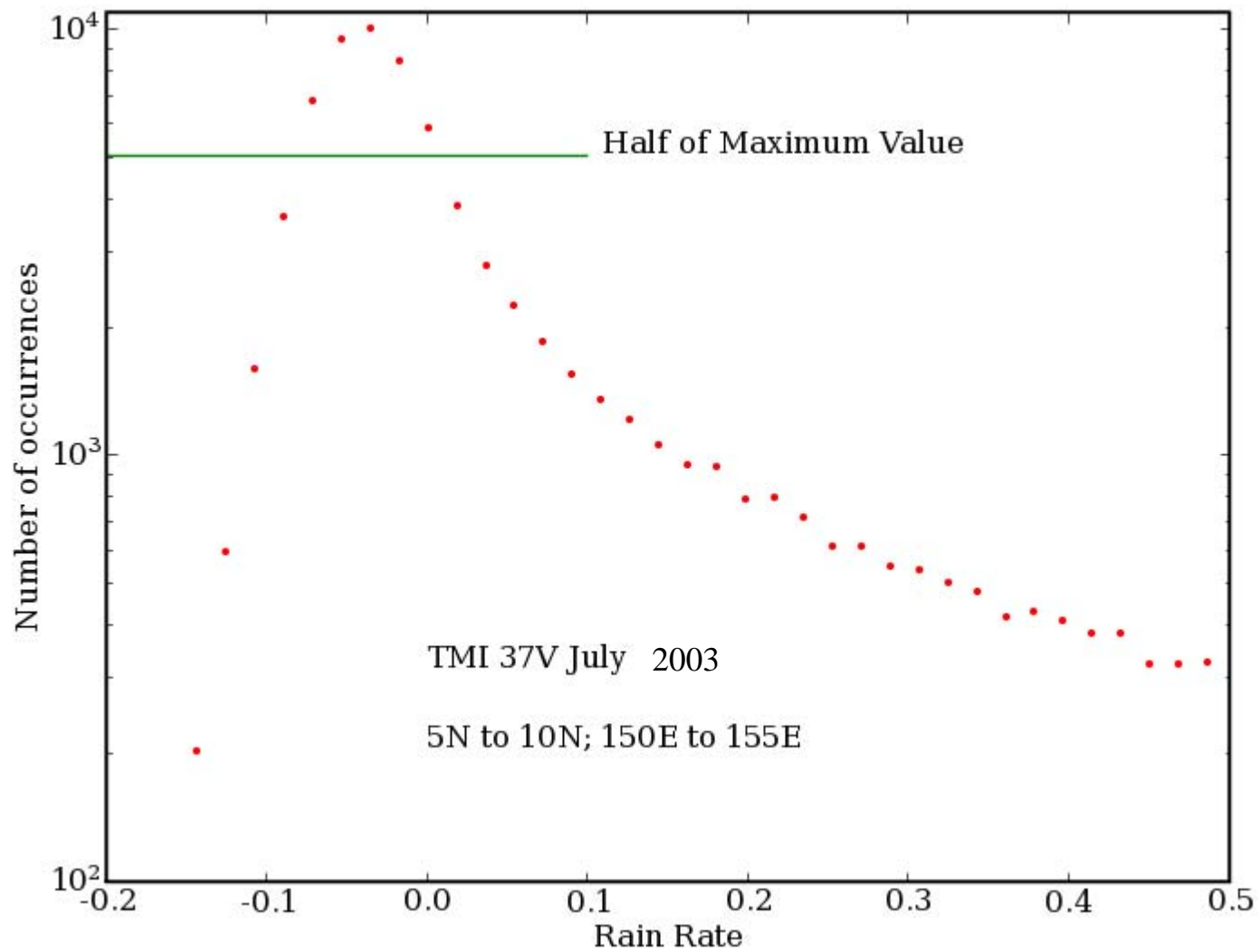
Department of Atmospheric Science

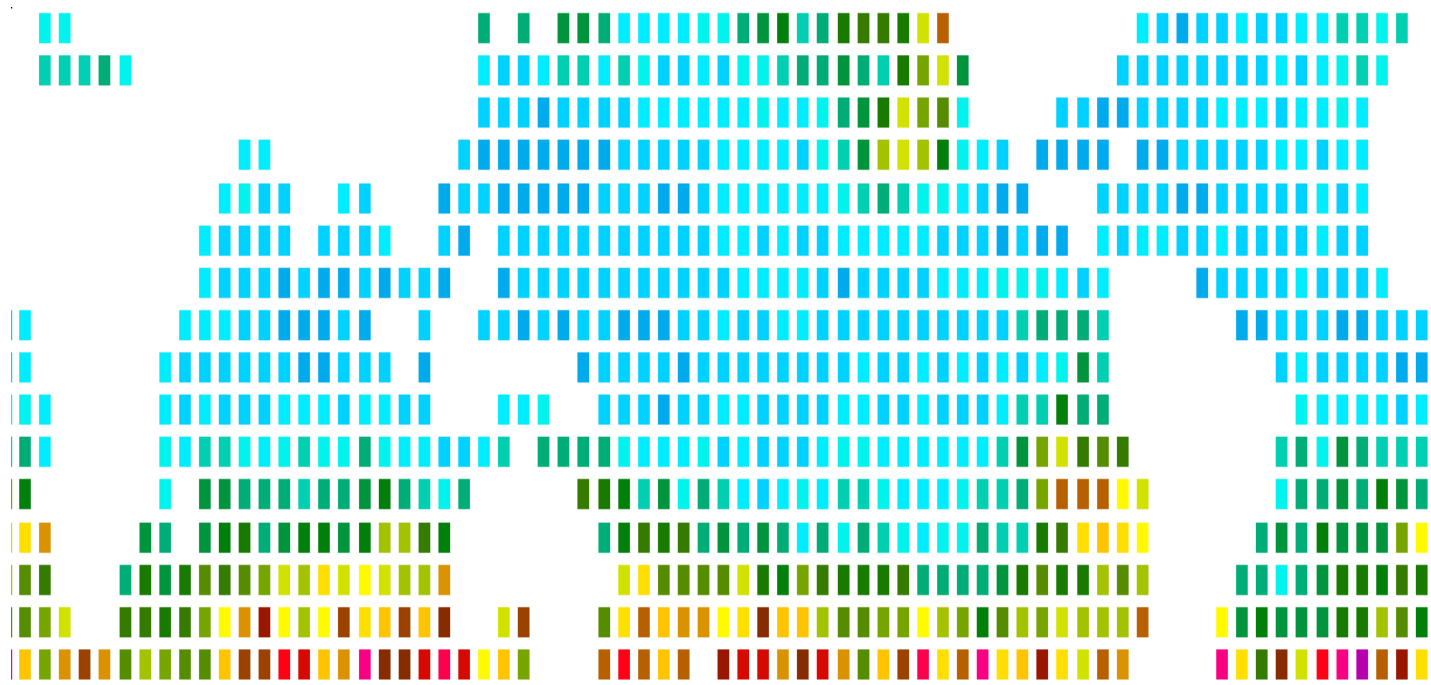
Texas A&M University

College Station, TX; USA

Two Pass Level 3 Oceanic Rainfall Algorithm

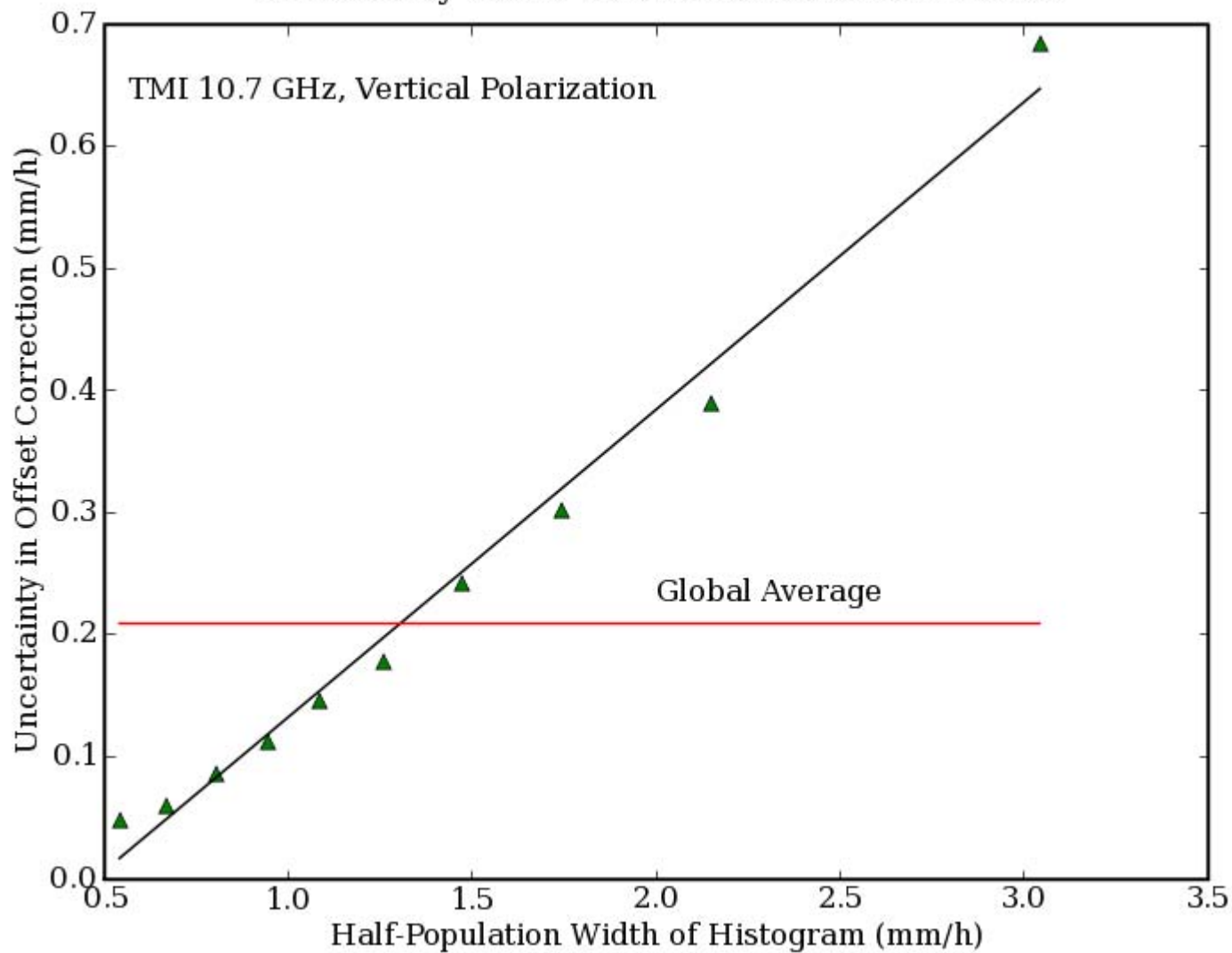






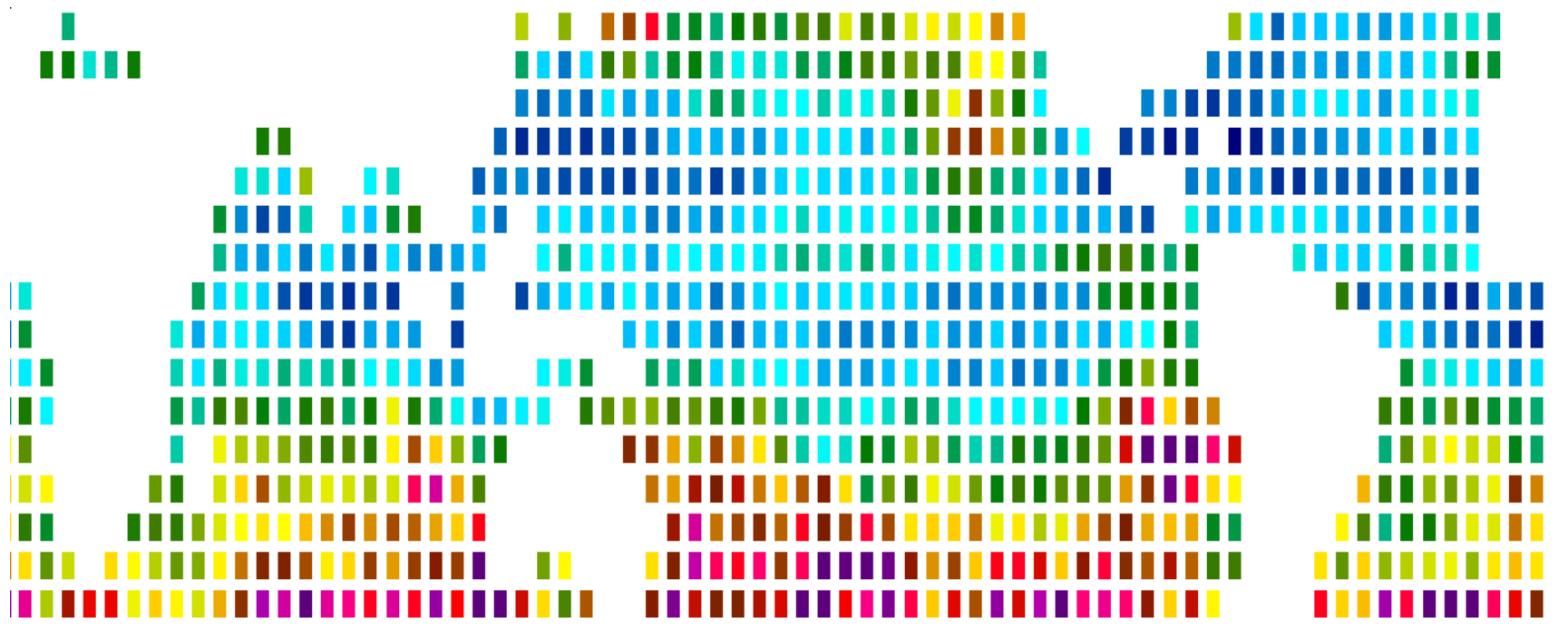
FULL WIDTH OF HALF POPULATION OF HISTOGRAMS JULY 2003
 TMI 37V CHANNEL

Uncertainty in the Rain Rate Offset Correction





OFFSET CONTRIBUTION TO RAIN UNCERTAINTY, TMI JULY 2003
CONSTANT OFFSET UNCERTAINTIES



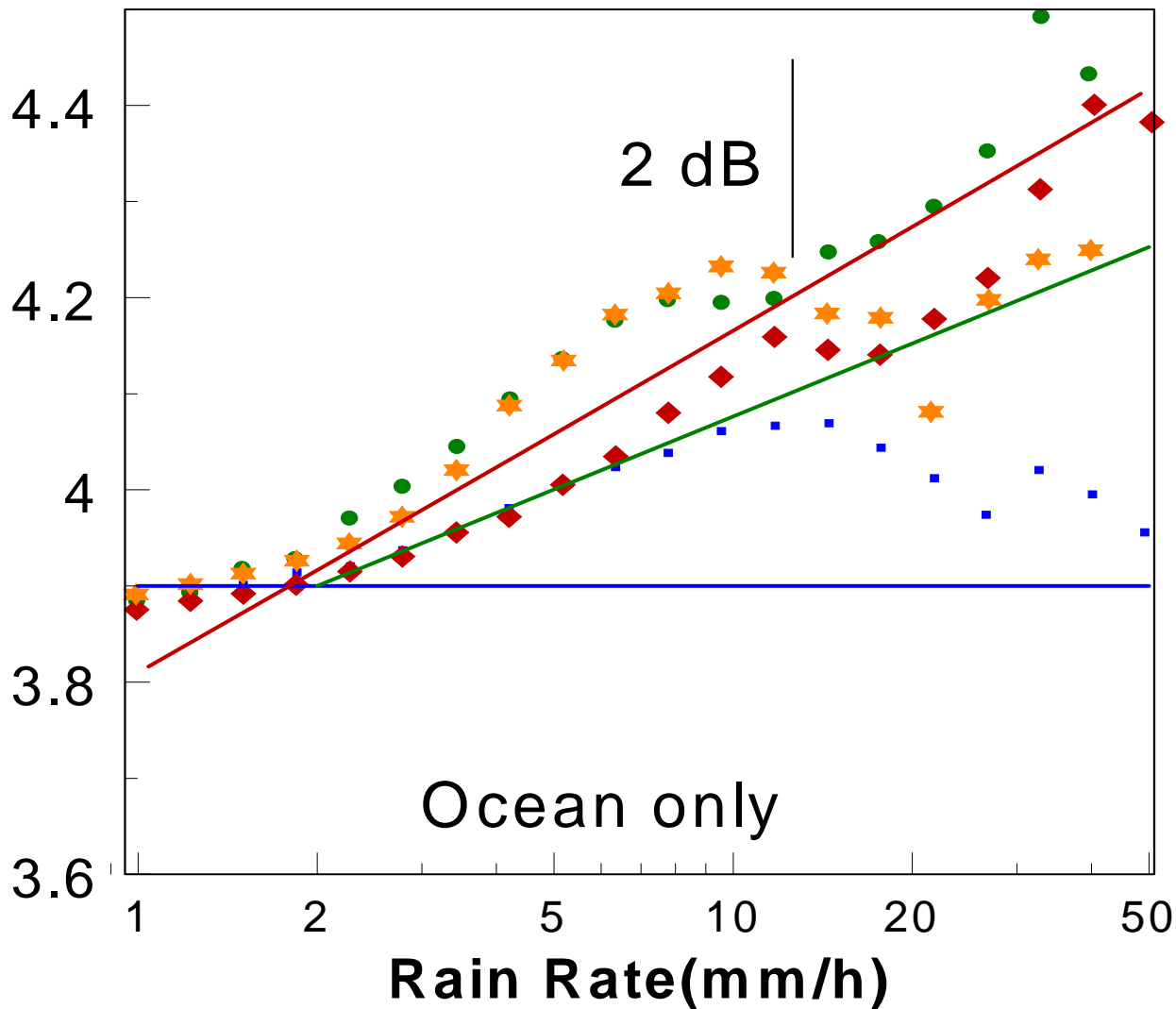
0.00

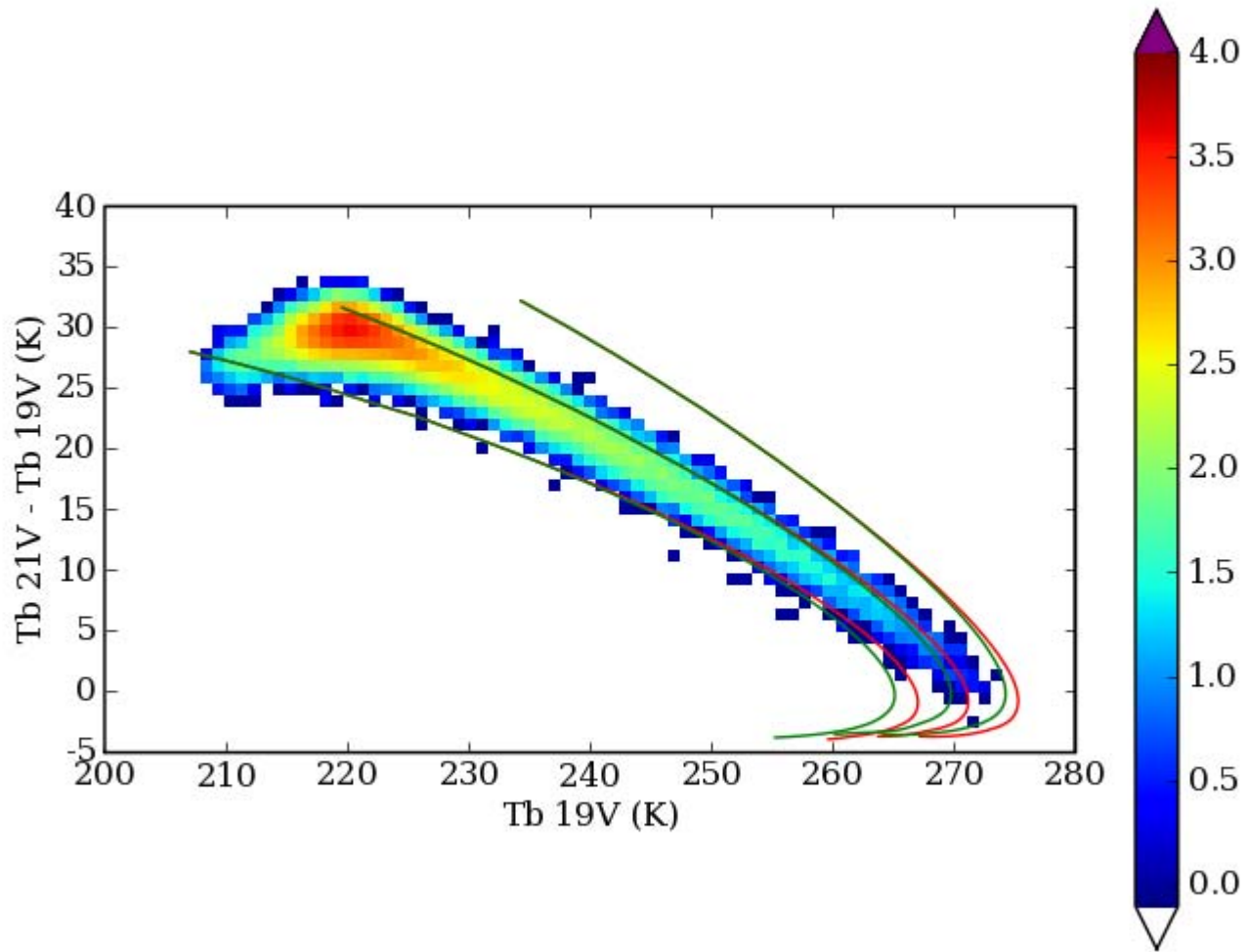
mm/day

2.5

OFFSET CONTRIBUTION TO RAIN UNCERTAINTY, TMI JULY 2003
VARIABLE OFFSET UNCERTAINTIES

Log10(N0)





TRMM Precipitation Radar Data for 1998

Profiles expressed as Surface Rain Rate & Slope

Surface Rain Is Desired Quantity but Column Mean is Closer to What the Radiometer Measures

Two Ways of Computing TBs

1) Compute TBs at Radar Resolution and Smooth to Radiometer Res

Approximates Actual Measurement

1) Smooth Rain Field to Radiometer Resolution Then Compute TBs

Implicit in Interpretation

Don't Know Structure in General Case

The Difference is the Beam Filling Error

Vertical Profile Error Remains

Can Only Be Corrected in a Statistical Sense

How can we make sense of this mess?

Expand $T_b(RR)$ around $\langle RR \rangle$ (mean RR over FOV)

$$\Delta T_b = \frac{1}{2} \frac{\partial^2 T_b}{\partial RR^2} \sigma_{RR}^2$$

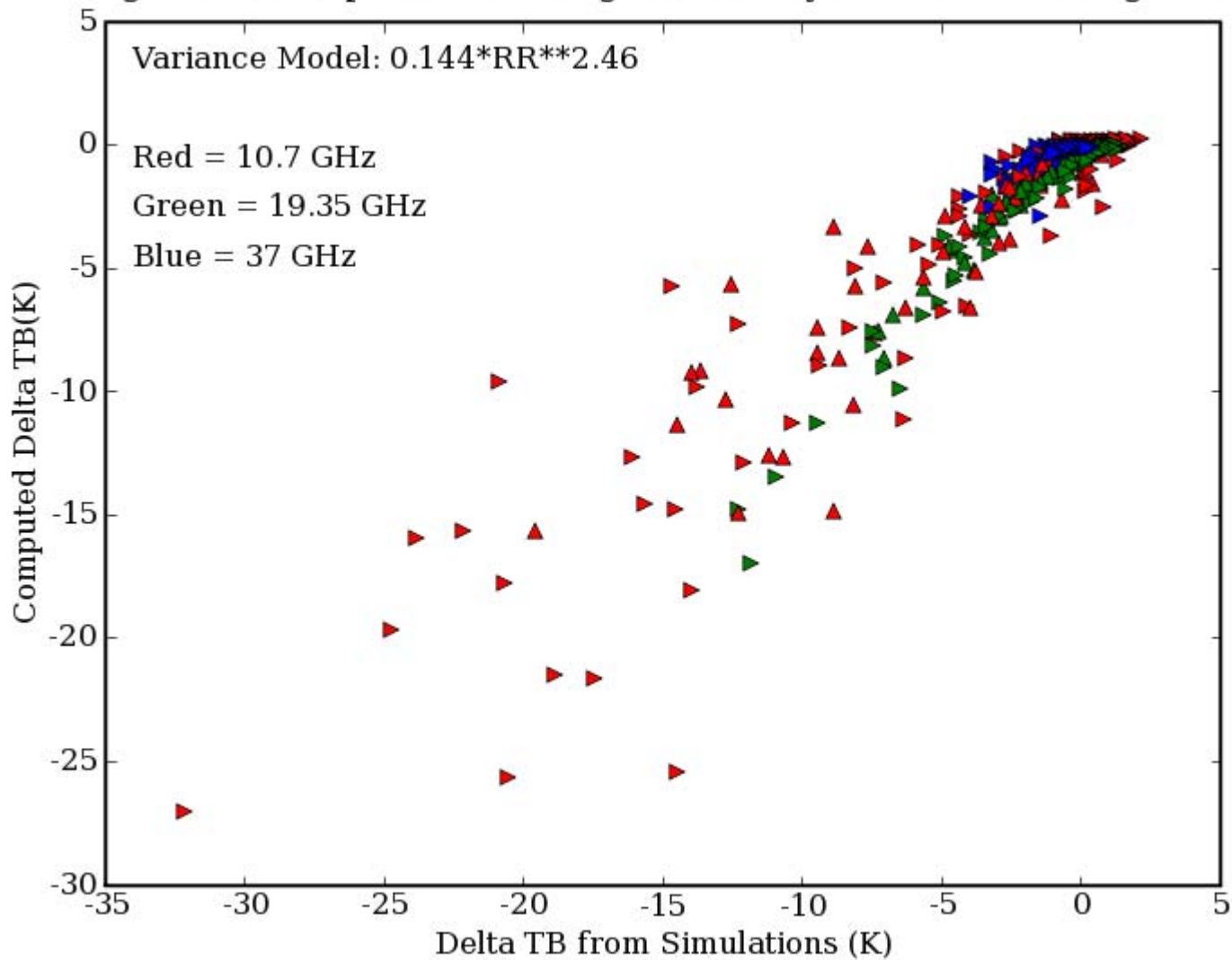
To a Very Good Approximation:

$$T_b = T_0 + (280K - T_0) * (1 - \exp(-RR / R_c)) + A * RR$$

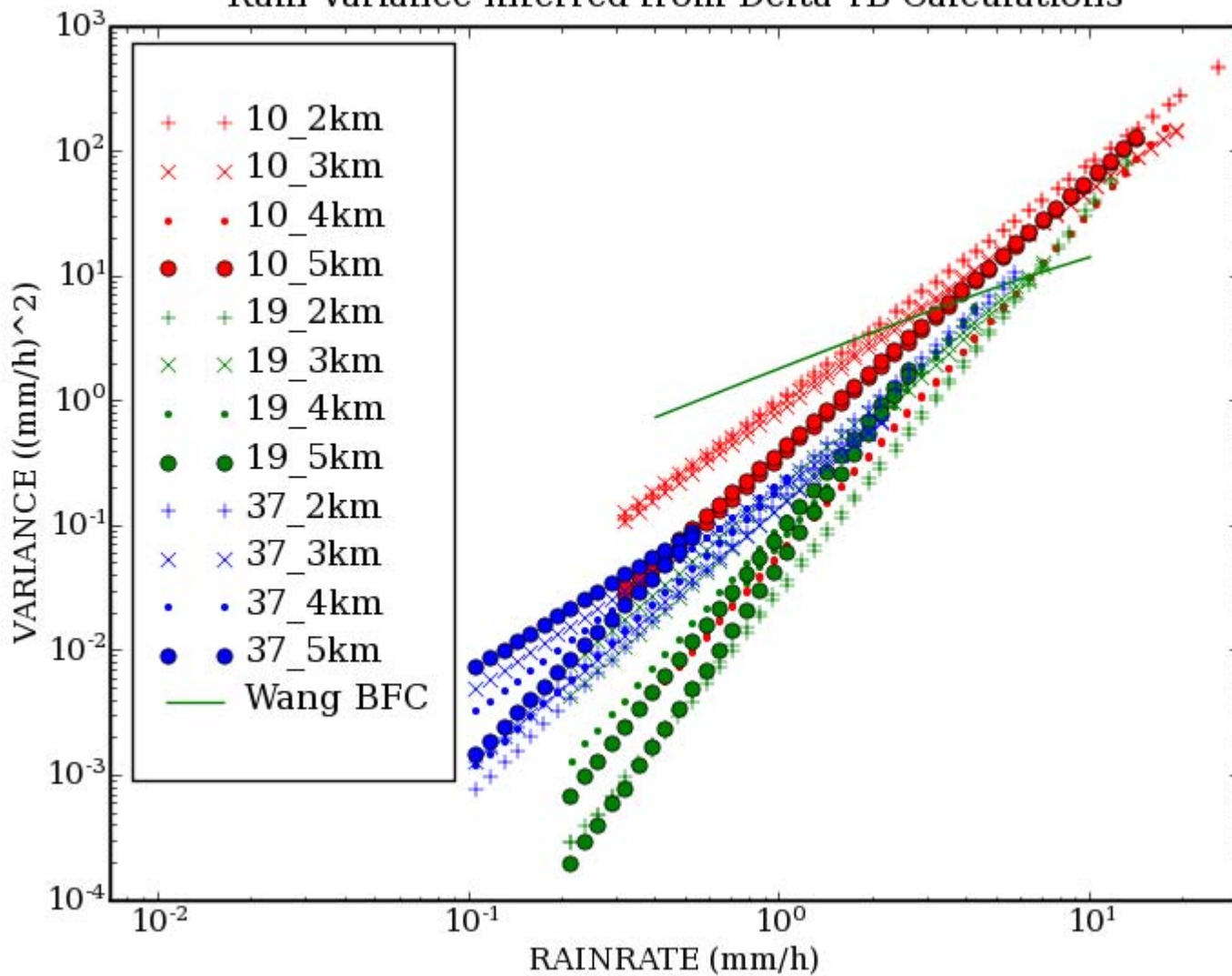
R_c , T_0 and A are Functions of Freq., Pol. and Freezing Level

Thus, ΔT_b Can be Expressed as σ_{RR}^2 over FOV

Brightness Temperature Change caused by Rainfall Inhomogeneity



Rain Variance Inferred from Delta TB Calculations



Conclusions

Improved description of Offset Uncertainty

Reduces Net Uncertainty in Most Locations

Improved Drop Size Distribution

Slight Reduction of Uncertainty but firmer basis

Improved Beam Filling Correction

In Progress